

On the source of orogenic gold

Andrew G. Tomkins

School of Geosciences, Monash University, Melbourne, VIC 3800, Australia

Gold has historically been a key strategic commodity, and the study of gold deposit formation has long been investigated. Debate on the genesis of so-called “orogenic gold deposits” (>75% of gold recovered through history; Phillips, 2013), continues largely because it is difficult to reliably identify the source of gold. In this issue of *Geology*, Gaboury (2013, p. 1207) identifies a geochemical tracer in orogenic gold deposits, ethane (C_2H_6), which must be sourced from carbonaceous metasedimentary rocks.

Orogenic gold deposits (Böhlke, 1982) dominantly form in metamorphic rocks in the mid- to shallow crust (5–15 km depth), at or above the brittle-ductile transition, in compressional settings that facilitate transfer of hot gold-bearing fluids from deeper levels (Goldfarb et al., 2005; Groves et al., 1998; Phillips and Powell, 2009). The term “orogenic” is used because these deposits likely form in accretionary and collisional orogens (Groves et al., 1998). Transfer of weakly oxidized, low-salinity fluids to the sites of gold deposition is controlled by earthquake events (Cox, 2005; Sibson et al., 1988), allowing fluids to rapidly traverse large thicknesses of crust. This rapid rise takes the fluids out of equilibrium with their surroundings, promoting destabilization of the gold-carrying hydrosulfide complexes $[Au(HS)_2^-]$ and $AuHS$. The chemical cause of gold precipitation, facilitated by a temperature-pressure decrease, varies from place to place, and mechanisms such as fluid-rock reaction (Evans et al., 2006), boiling (Weatherley and Henley, 2013), fluid mixing (Bateman and Hagemann, 2004), and chemisorption on surfaces of pyrite and arsenopyrite (e.g., Möller and Kersten, 1994) have been proposed. We can investigate the cause of gold deposition by sampling rocks where precipitation occurred, and structures through which fluids flowed. In contrast, the question where the gold comes from is more difficult, because the potential sources are distant to the sampled gold deposits.

There are two plausible sources for the gold: (1) metamorphic rocks, from which fluids are generated as temperatures increase; and (2) felsic-intermediate magmas, which release fluids as they crystallize. Gold-bearing magmatic-hydrothermal deposits are enriched in many elements, including S, Cu, Mo, Sb, Bi, W, Pb, Zn, Te, Hg, As, and Ag (e.g., Goldfarb et al., 2005; Richards, 2009). Such deposits have been referred to as gold-plus deposits (e.g., Phillips, 2013), but most orogenic gold deposits fall into the alternative group of gold-only deposits, and are more enigmatic. These are characterized by elevated S and As, and have only minor enrichments in the other elements. The current dominant opinion is that metamorphic rocks are the source for these deposits (Goldfarb et al., 2005; Phillips and Powell, 2010).

In the early 1990s, the Crustal Continuum Model (Groves, 1993) was developed to explain the occurrence of orogenic gold deposits in metamorphic rocks of granulite to sub-greenschist facies. It was suggested that gold deposits can form at almost any depth in the crust, at temperatures from 750 °C to 180 °C. More recently, it was shown that large fluid volumes cannot migrate effectively through rocks hotter than ~650 °C (Tomkins and Grundy, 2009; Phillips and Powell, 2009). Then, the continuum was further constrained by showing that the most fertile metamorphic fluids are released over ~500–550 °C, with little fluid released beyond ~550 °C (Tomkins, 2010). These constraints are consistent with observations that most orogenic gold deposits occur in greenschist to lower amphibolite facies rocks (Groves, 1993).

Xue et al. (2013), however, used the $\Delta^{33}S$ signature of sulfur in four Archean orogenic gold deposits to suggest a granitic source. The same result was independently obtained for different Archean deposits using the signature of noble gases and halogens in fluids released by crushing of gold-associated quartz-carbonate veins (Kendrick et al., 2011). In some Phanerozoic slate belts, however, there can be little debate about a metamorphic source. For example, metamorphism of turbidites in New Zealand’s Southern Alps caused Au, S, and As release from diagenetic pyrite, allowing gold deposit formation in the absence of intrusions (Pitcairn et al., 2006). Similarly, in the more globally significant Victorian goldfields (Australia), there are no intrusions coeval with the largest gold deposits (e.g., Wilson et al., 2013), whereas a 35-km-thick sequence of fertile metamorphic rocks underlies the deposits.

The vast majority of orogenic gold (excluding Witwatersrand, South Africa) is from three periods in geologic time: the Neoproterozoic (ca. 2700–2400 Ma), a second period in the Paleoproterozoic (ca. 2100–1800 Ma), and a third period from ca. 650 Ma continuing throughout the Phanerozoic (Goldfarb et al., 2001). Two explanations have been offered for this timing: (1) because orogenic gold deposit formation requires accretionary tectonics, the major periods of formation coincided with periods of continental growth (Goldfarb et al., 2001), and (2) during the Phanerozoic, increased ocean oxygenation facilitated uptake of gold in biogenic and diagenetic pyrite, which became the gold source during later accretion and metamorphism (Tomkins, 2013). The first explanation must be correct to some extent, but cannot explain the relative lack of gold during the formation of Rodinia; the second requires that gold can be sourced from carbonaceous metasedimentary rocks.

Some proponents of the metamorphic source model suggest that hydrated mafic rocks are the source for fluids and gold (Powell et al., 1991; Phillips and Powell, 2009, 2010; Wilson et al., 2013), but thermodynamic calculations suggest that metamorphism of pyritic carbonaceous sedimentary rocks can generate much more fertile fluids than mafic rocks (Tomkins, 2010). For mafic and metasedimentary rocks, the most fertile fluids are generated during high-temperature, low-pressure metamorphism, constraining the tectonic settings where gold deposits are likely to form. Inverted backarc basins should be the ideal tectonic setting, because heat propagating through the crust as a consequence of rifting continues throughout the period of backarc inversion (heat flow is much slower than tectonic plate movements). This continued heat through-flow drives widespread high-temperature, low-pressure metamorphism of the ideal source rocks during the ideal structural period, allowing contemporaneous gold concentration across the very large regions where these deposits are found (Tomkins, 2010).

Independent suggestions that carbonaceous metasedimentary rocks are an ideal source came when development of laser ablation mapping allowed recognition that biogenic and diagenetic pyrite grains contain elevated Au, As, and other elements enriched in orogenic gold deposits (Large et al., 2011, 2009; Thomas et al., 2011). In contrast, mafic rocks are comparatively deficient in As, Ag, Pb, Zn, and Sb.

Gaboury (2013) makes two important points: (1) C_2H_6 was found in Detour Lake, Canada’s largest gold deposit, where there are no known adjacent carbonaceous metasedimentary rocks, and (2) this orogenic deposit is hosted in upper greenschist facies metamorphic rocks, relatively close to the greenschist-amphibolite transition where metamorphic fluids

*Email: andy.tomkins@monash.edu.

are generated. Point 2 implies a short fluid pathway from the site of liberation to the site of gold deposition, thus lessening the chance of contribution from multiple sources, making it probable that carbonaceous metasedimentary rocks were the source. Gaboury also explains the occurrence of CO₂-rich-H₂O-poor fluids at many large gold deposits through the reaction of C₂H₆ with H₂O to form CO₂, implying that carbonaceous metasedimentary source rocks are globally important.

Many questions remain about the source of orogenic gold. Some form of the continuum model may remain valid, with a temperature range of ~550–300 °C (cf. Phillips and Powell, 2009), involving formation of deposits from the deeper levels of the brittle-ductile transition to the shallow reaches of earthquake failure planes. The Archean and Paleoproterozoic orogenic gold-forming systems appear to be subtly different from the Phanerozoic ones, in terms of tectonics, source rocks, and deposit settings; thus, we need to be careful about extrapolating from one group to the other. One important unanswered question is the time needed to form orogenic gold deposits. If they form quickly, igneous sources might be more probable, but if they form over extended periods, metamorphic sources would be more probable. In terms of metamorphic sources, Gaboury's (2013) work adds weight to the evidence suggesting that carbonaceous metasedimentary rocks appear chemically more favorable, though mafic rocks remain a plausible source.

REFERENCES CITED

- Bateman, R., and Hagemann, S., 2004, Gold mineralisation throughout about 45 Ma of Archean orogenesis: Protracted flux of gold in the Golden Mile, Yilgarn craton, Western Australia: *Mineralium Deposita*, v. 39, p. 536–559, doi:10.1007/s00126-004-0431-2.
- Böhlke, J.K., 1982, Orogenic (metamorphic-hosted) gold-quartz veins.: U.S. Geological Survey Open-file Report, v. 82–795, p. 70–76.
- Cox, S.F., 2005, Coupling between deformation, fluid pressures, and fluid flow in ore-producing hydrothermal systems at depth in the crust, *in* Hedenquist, J. W., Thompson, J.F.H., Goldfarb, R.J., and Richards, J.P., eds., *Economic Geology 100th Anniversary Volume 1905–2005*: Littleton, Colorado, Society of Economic Geologists, p. 39–75.
- Evans, K.A., Phillips, G.N., and Powell, R., 2006, Rock-buffering of auriferous fluids in altered rocks associated with the Golden Mile-style mineralization, Kalgoorlie gold field, Western Australia: *Economic Geology and the Bulletin of the Society of Economic Geologists*, v. 101, p. 805–817, doi:10.2113/gsecongeo.101.4.805.
- Gaboury, D., 2013, Does gold in orogenic deposits come from pyrite in deeply buried carbon-rich sediments?: Insight from volatiles in fluid inclusions: *Geology*, v. 41, p. 1207–1210, doi:10.1130/G34788.1.
- Goldfarb, R.J., Baker, T., Dube, B., Groves, D.I., Hart, C.J.R., and Gosselin, P., 2005, Distribution, character, and genesis of gold deposits in metamorphic terranes, *in* Hedenquist, J. W., Thompson, J. F. H., Goldfarb, R. J., and Richards, J. P., eds., *Economic Geology. 100th Anniversary Volume 1905–2005*: Littleton, Colorado, Society of Economic Geologists, p. 407–450.
- Goldfarb, R.J., Groves, D.I., and Gardoll, S., 2001, Orogenic gold and geologic time: a global synthesis: *Ore Geology Reviews*, v. 18, p. 1–75, doi:10.1016/S0169-1368(01)00016-6.
- Groves, D.I., 1993, The crustal continuum model for late-Archaean lode-gold deposits of the Yilgarn Block, Western Australia: *Mineralium Deposita*, v. 28, p. 366–374, doi:10.1007/BF02431596.
- Groves, D.I., Goldfarb, R.J., Gebre-Mariam, M., Hagemann, S.G., and Robert, F., 1998, Orogenic gold deposits: a proposed classification in the context of their crustal distribution and relationship to other gold deposit types: *Ore Geology Reviews*, v. 13, p. 7–27, doi:10.1016/S0169-1368(97)00012-7.
- Kendrick, M.A., Honda, M., Walshe, J., and Petersen, K., 2011, Fluid sources and the role of abiogenic-CH₄ in Archean gold mineralization: Constraints from noble gases and halogens: *Precambrian Research*, v. 189, p. 313–327, doi:10.1016/j.precamres.2011.07.015.
- Large, R.R., Bull, S.W., and Maslennikov, V., 2011, A carbonaceous sedimentary source-rock model for Carlin-type and orogenic gold deposits: *Economic Geology and the Bulletin of the Society of Economic Geologists*, v. 106, p. 331–358, doi:10.2113/econgeo.106.3.331.
- Large, R.R., Danyushevsky, L., Hollit, C., Maslennikov, V., Meffre, S., Gilbert, S., Bull, S., Scott, R., Emsbo, P., Thomas, H., Singh, B., and Foster, J., 2009, Gold and trace element zonation in pyrite using a laser imaging technique: Implications for the timing of gold in orogenic and carlin-style sediment-hosted deposits: *Economic Geology and the Bulletin of the Society of Economic Geologists*, v. 104, p. 635–668, doi:10.2113/gsecongeo.104.5.635.
- Möller, P., and Kersten, G., 1994, Electrochemical accumulation of visible gold on pyrite and arsenopyrite surfaces: *Mineralium Deposita*, v. 29, p. 404–413, doi:10.1007/BF01886958.
- Phillips, G.N., Australian and global setting for gold in 2013, *in* *Proceedings World Gold 2013*, Brisbane, Australia, 26–29 September, 2013: The Australian Institute of Mining and Metallurgy, p. 15–21.
- Phillips, G.N., and Powell, R., 2009, Formation of gold deposits: Review and evaluation of the continuum model: *Earth-Science Reviews*, v. 94, p. 1–21, doi:10.1016/j.earscirev.2009.02.002.
- Phillips, G.N., and Powell, R., 2010, Formation of gold deposits: A metamorphic devolatilization model: *Journal of Metamorphic Geology*, v. 28, p. 689–718, doi:10.1111/j.1525-1314.2010.00887.x.
- Pitcairn, I.K., Teagle, D.A.H., Craw, D., Olivo, G.R., Kerrich, R., and Brewer, T.S., 2006, Sources of metals and fluids in orogenic gold deposits: Insights from the Otago and Alpine Schists, New Zealand: *Economic Geology and the Bulletin of the Society of Economic Geologists*, v. 101, p. 1525–1546, doi:10.2113/gsecongeo.101.8.1525.
- Powell, R., Will, T.M., and Phillips, G.N., 1991, Metamorphism in Archaean greenstone belts: Calculated fluid compositions and implications for gold mineralisation: *Journal of Metamorphic Geology*, v. 9, p. 141–150, doi:10.1111/j.1525-1314.1991.tb00510.x.
- Richards, J.P., 2009, Postsubduction porphyry Cu-Au and epithermal Au deposits: Products of remelting of subduction-modified lithosphere: *Geology*, v. 37, p. 247–250, doi:10.1130/G25451A.1.
- Sibson, R.H., Robert, F., and Poulsen, K.H., 1988, High-angle reverse faults, fluid-pressure cycling, and mesothermal gold-quartz deposits: *Geology*, v. 16, p. 551–555, doi:10.1130/0091-7613(1988)016<0551:HARFFP>2.3.CO;2.
- Thomas, H.V., Large, R.R., Bull, S.W., Maslennikov, V., Berry, R.F., Fraser, R., Froud, S., and Moye, R., 2011, Pyrite and pyrrhotite textures and composition in sediments, laminated quartz veins, and reefs at Bendigo gold mine, Australia: Insights for ore genesis: *Economic Geology and the Bulletin of the Society of Economic Geologists*, v. 106, p. 1–31, doi:10.2113/econgeo.106.1.1.
- Tomkins, A.G., 2010, Windows of metamorphic sulfur liberation in the crust: Implications for gold deposit genesis: *Geochimica et Cosmochimica Acta*, v. 74, p. 3246–3259, doi:10.1016/j.gca.2010.03.003.
- Tomkins, A.G., 2013, A biogeochemical influence on the secular distribution of orogenic gold: *Economic Geology and the Bulletin of the Society of Economic Geologists*, v. 108, p. 193–197, doi:10.2113/econgeo.108.2.193.
- Tomkins, A.G., and Grundy, C., 2009, Upper Temperature Limits of Orogenic Gold Deposit Formation: Constraints from the Granulite-Hosted Griffin's Find Deposit, Yilgarn Craton: *Economic Geology and the Bulletin of the Society of Economic Geologists*, v. 104, p. 669–685, doi:10.2113/gsecongeo.104.5.669.
- Weatherley, D.K., and Henley, R.W., 2013, Flash vaporization during earthquakes evidenced by gold deposits: *Nature Geoscience*, v. 6, p. 294–298, doi:10.1038/ngeo1759.
- Wilson, C.J.L., Schaub, P., and Leader, L.D., 2013, Mineral precipitation in the quartz reefs of the Bendigo gold deposit, Victoria, Australia: *Economic Geology and the Bulletin of the Society of Economic Geologists*, v. 108, p. 259–278, doi:10.2113/econgeo.108.2.259.
- Xue, Y., Campbell, I.H., Ireland, T.R., Holden, P., and Armstrong, R., 2013, No mass-independent sulfur isotope fractionation in auriferous fluids supports a magmatic origin for Archean gold deposits: *Geology*, v. 41, p. 791–794, doi:10.1130/G34186.1.

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